

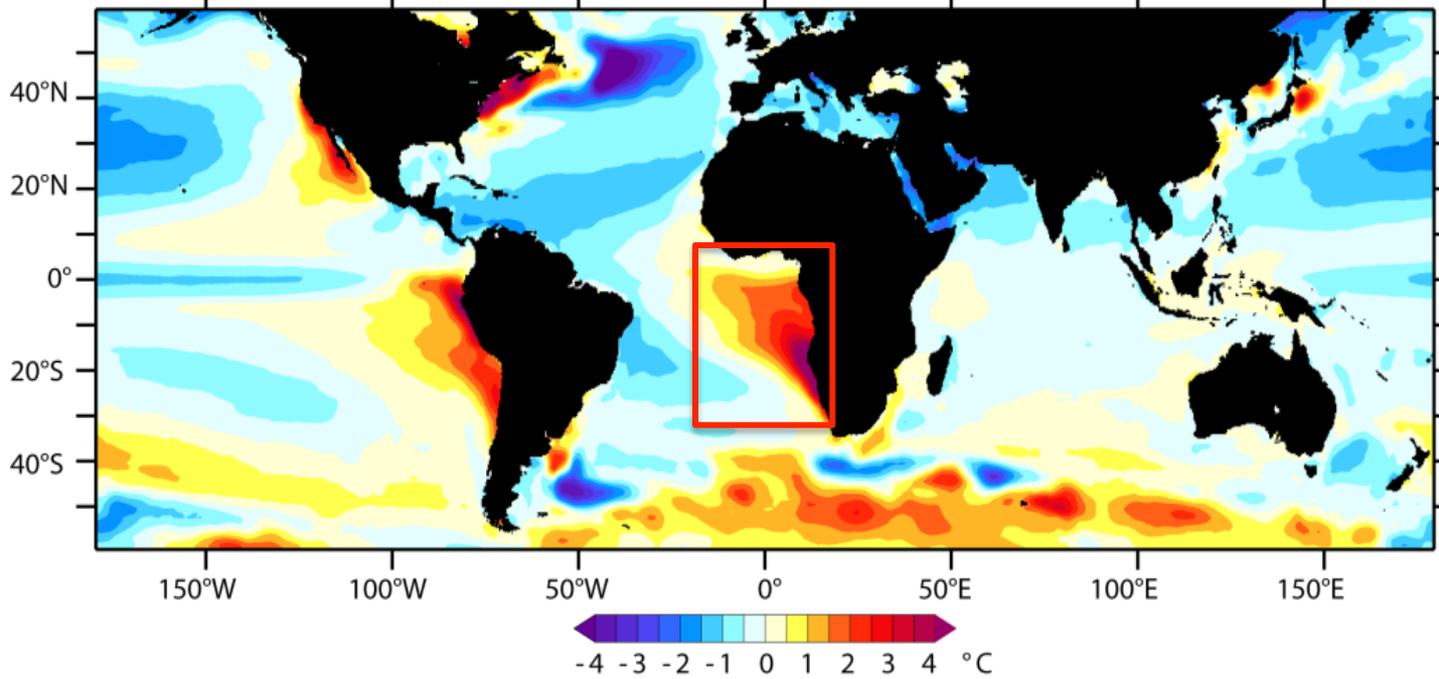
Role of equatorial forcing in SST bias development in the South-Eastern Tropical Atlantic in a high resolution version of CNRM-CM CGCM



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Mean SST bias in an ensemble of 39 models from CMIP5/historical

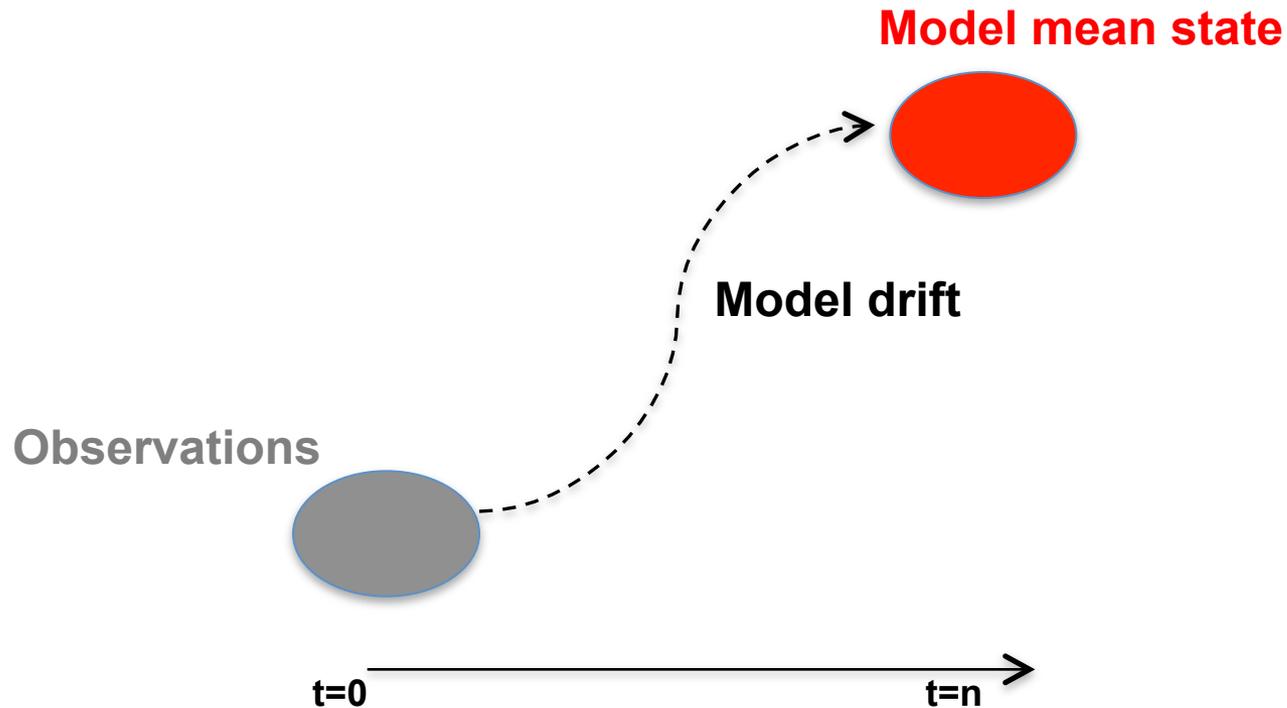


Progress and challenges:

- Systematic warm SST bias in S-E Tropical Atlantic (SETA)
- Quick bias development (weeks to months) (*full-field initialized seasonal hindcasts*)
- Different processes involved in SST errors: local/remote, atmosphere/ocean
- Model-dependent

Methodology:

analysis of model drift in seasonal hindcasts to study bias development



Goals:

1) What is the role of remote forcing from the Equator (westerly wind bias) in developing the warm SST error in the SETA ?

2) Is there an added value of higher resolution for simulating the SETA?

Model:

High and Low resolution version of CNRM-CM model

HR: ARPEGE (~50km) (T359L31) and NEMO ~0.25° (ORCA025, 75VL)

LR: ARPEGE (~140km) (T127L31) and NEMO ~1° (ORCA1, 42VL)

Experiments:

- ❑ **CTRL-HR / CTRL-LR: full field initialized seasonal hindcasts**
 - Initialization from ERAI and GLORYS2v3
 - Start date: 1 February
 - 10 years: 2000-2009
 - Three-members, 6 months lead time

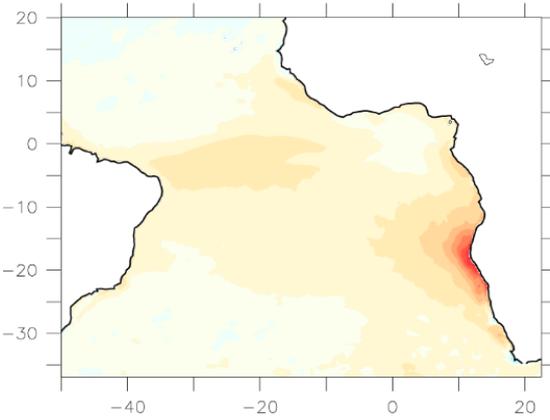
- ❑ **TAUEQ (HR):** wind stress from ERAI is replaced over the Equatorial Atlantic ($5^{\circ}\text{S} - 5^{\circ}\text{N}$)

1. CTRL-HR biases and their improvement in TAUEQ

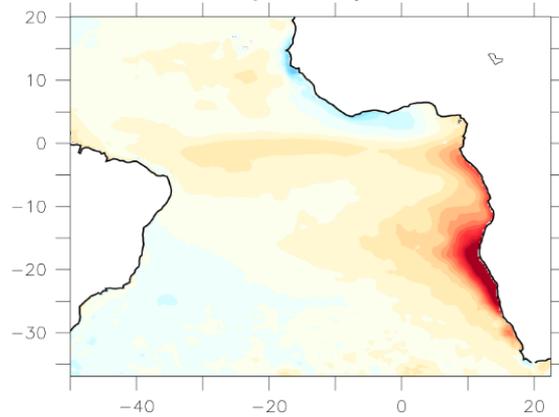
-- > CTRL-HR versus TAUEQ

SST bias evolution in CTRL-HR

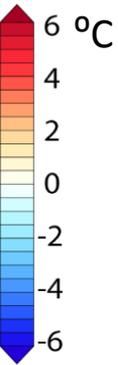
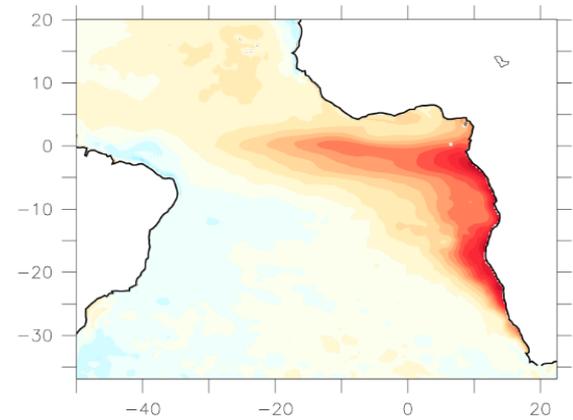
Feb - Mar



Apr - May



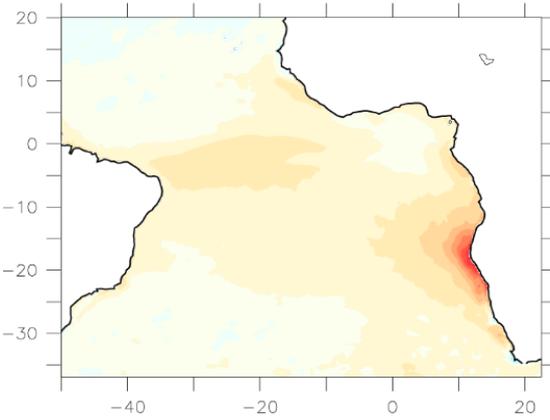
Jun - Jul



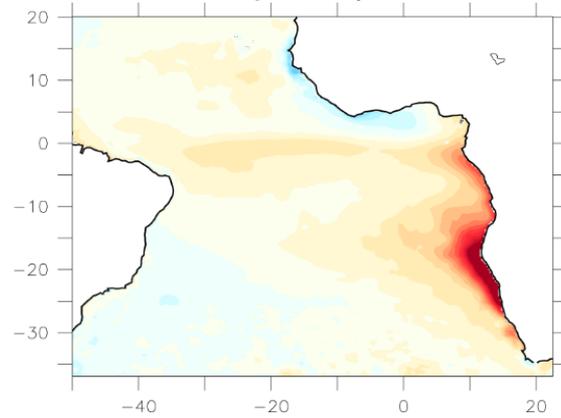
2-months averaged evolution of the SST bias with respect to GLORYS2v3

SST bias evolution in CTRL-HR

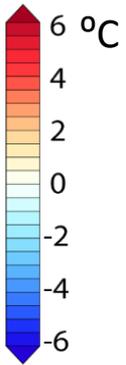
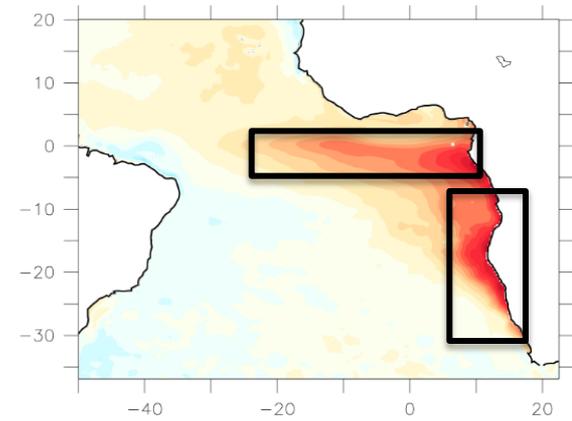
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Jun - Jul

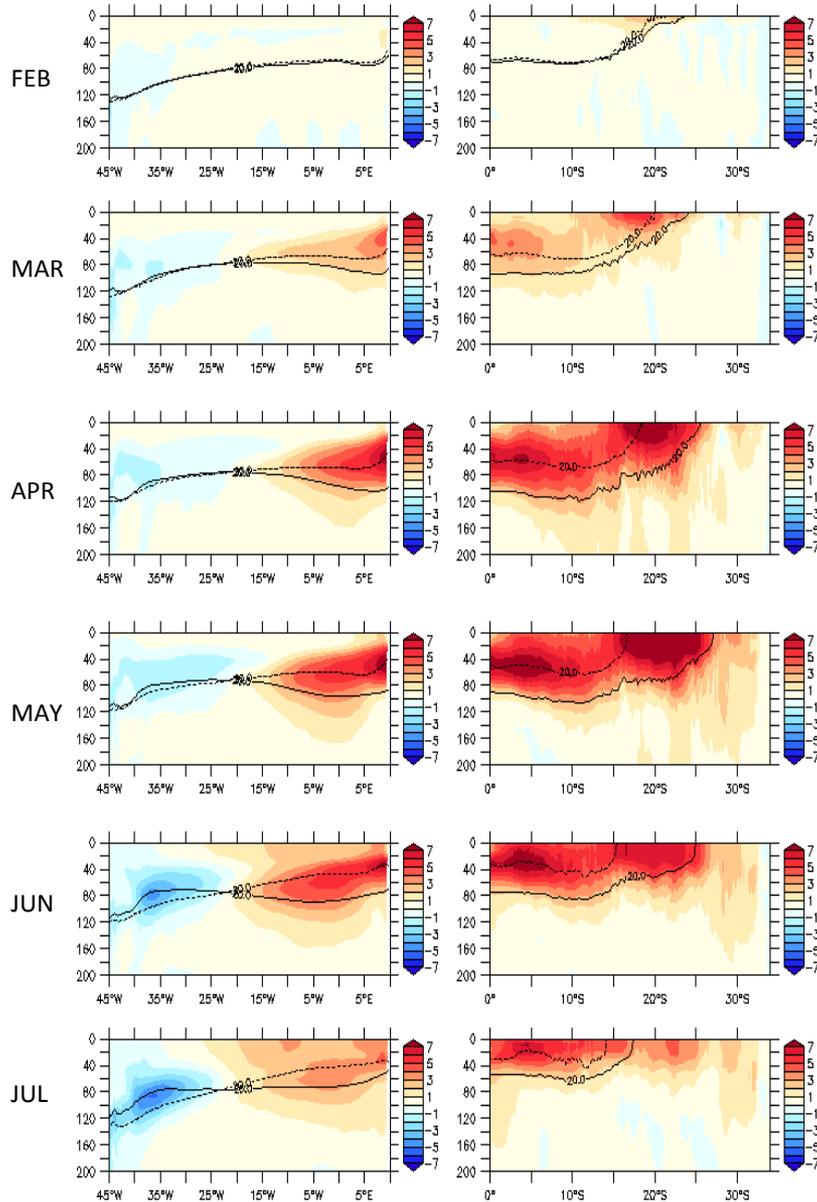


2-months averaged evolution of the SST bias with respect to GLORYS2v3

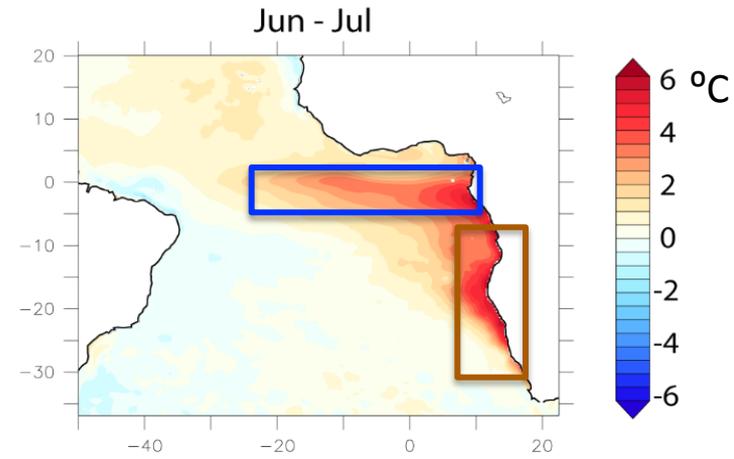
CTRL-HR

T along the equator (2°S-2°N)

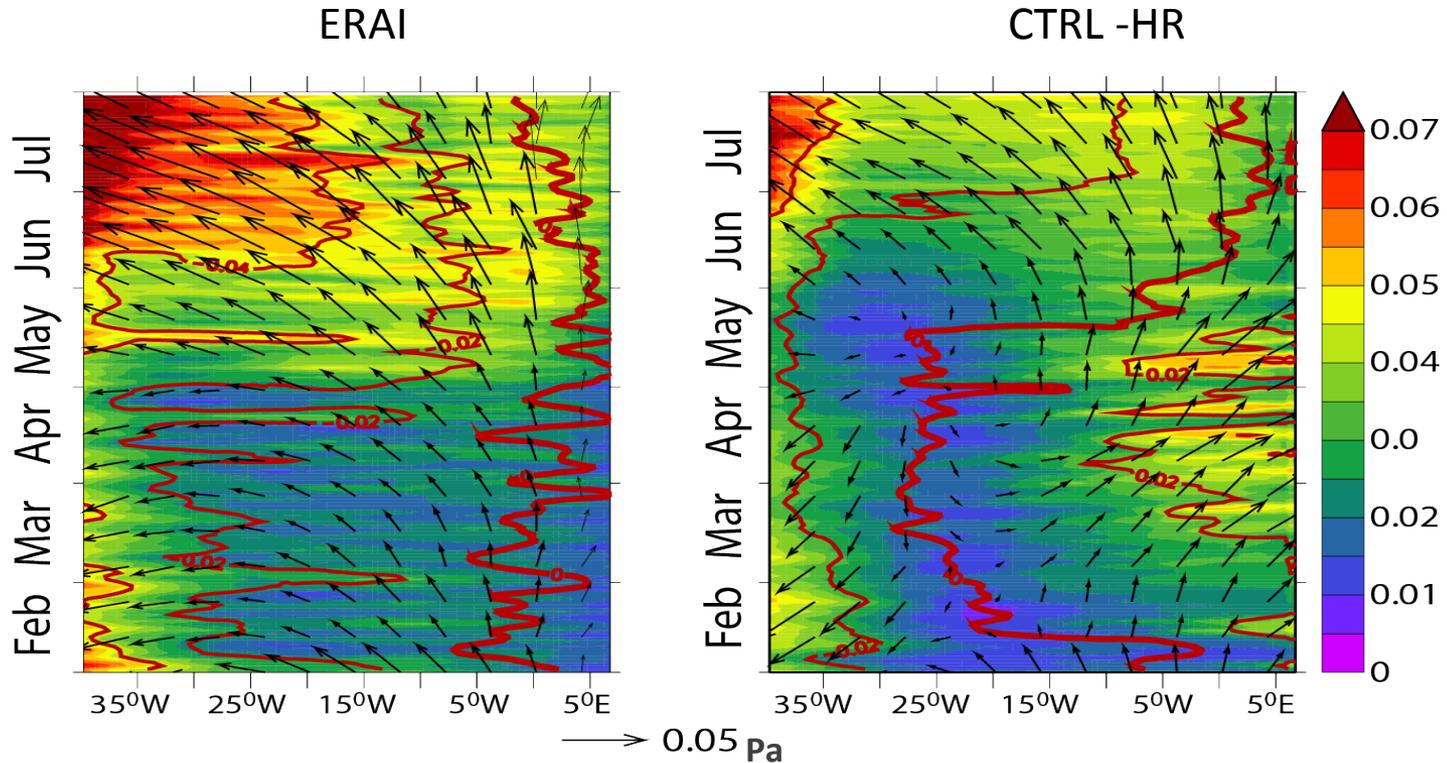
T along the coast (2°)



— Isotherme 20°C in the model
 - - - Isotherme 20°C in GLORYS2v3



Mean evolution of the equatorial wind stress (2°S- 2°N)



Shading : wind stress amplitude

Vector: wind stress vector

Contour: zonal wind stress

CTRL-HR

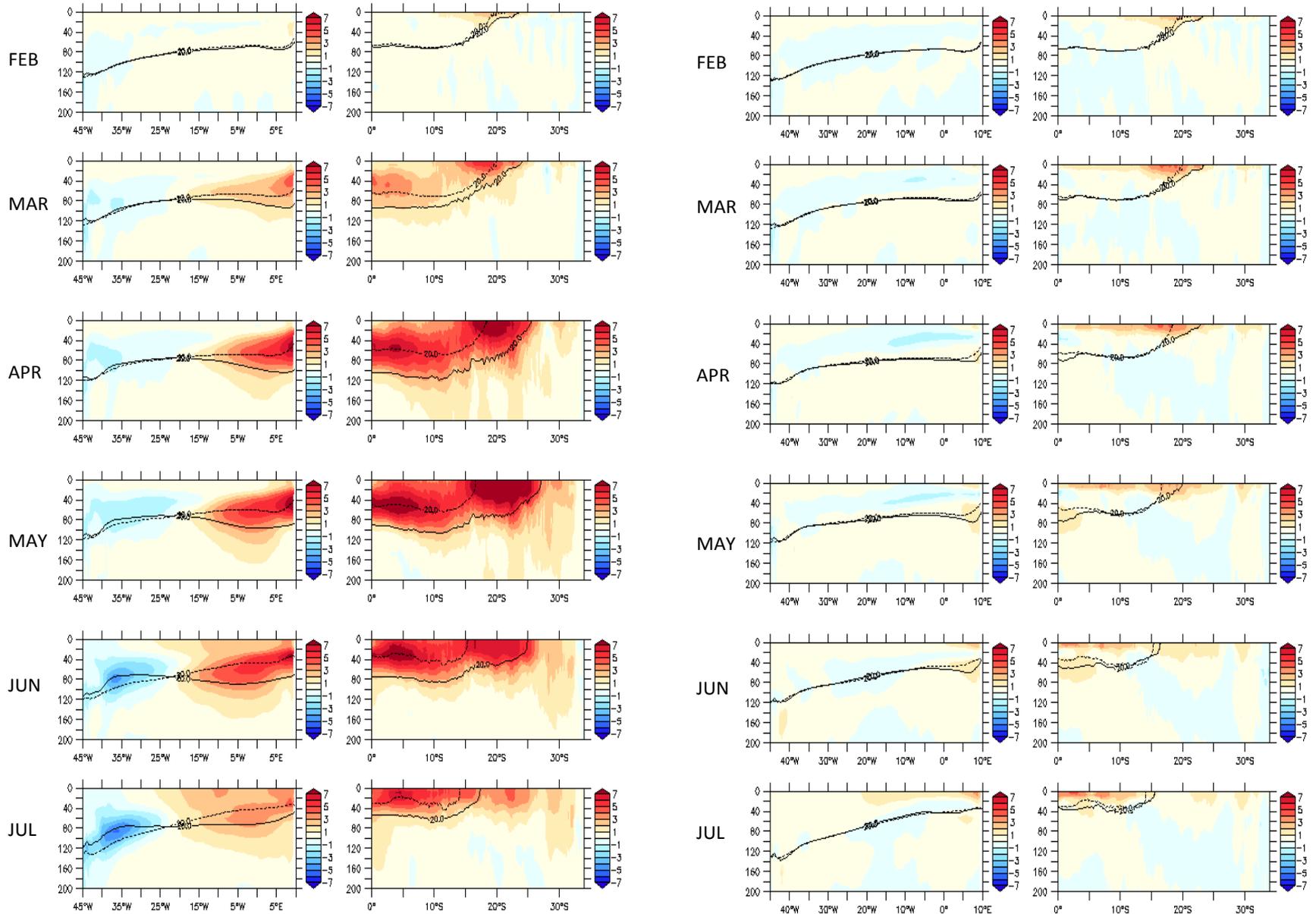
TAUEQ

T along the equator (2°S-2°N)

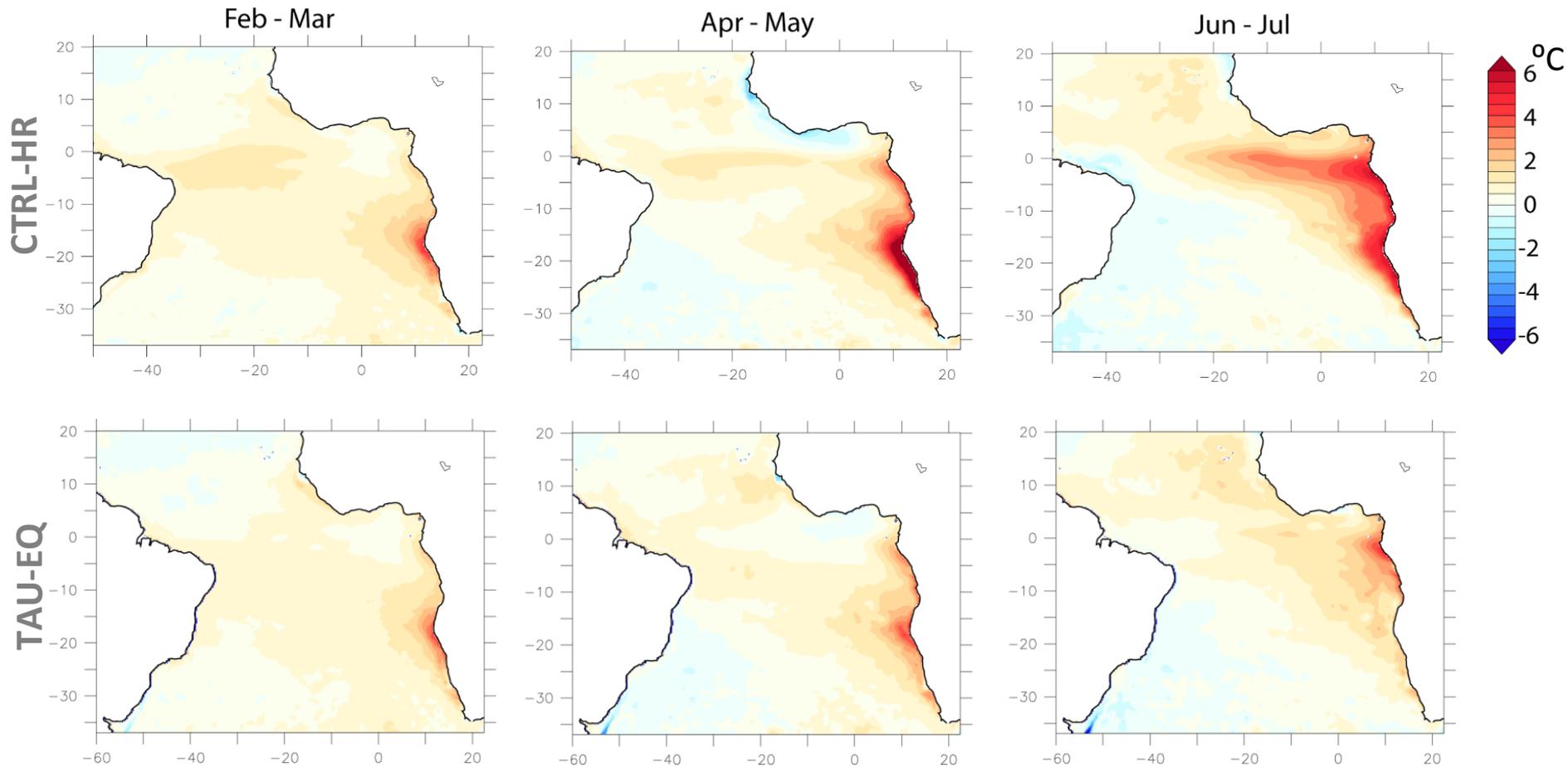
T along the coast (2°)

T along the equator (2°S-2°N)

T along the coast (2°)



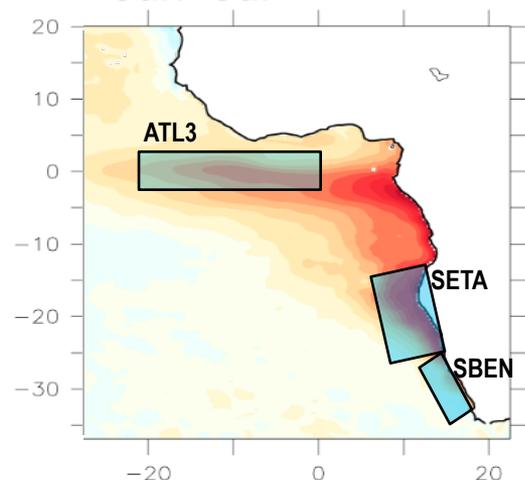
SST bias in respect to GLORYS2V3



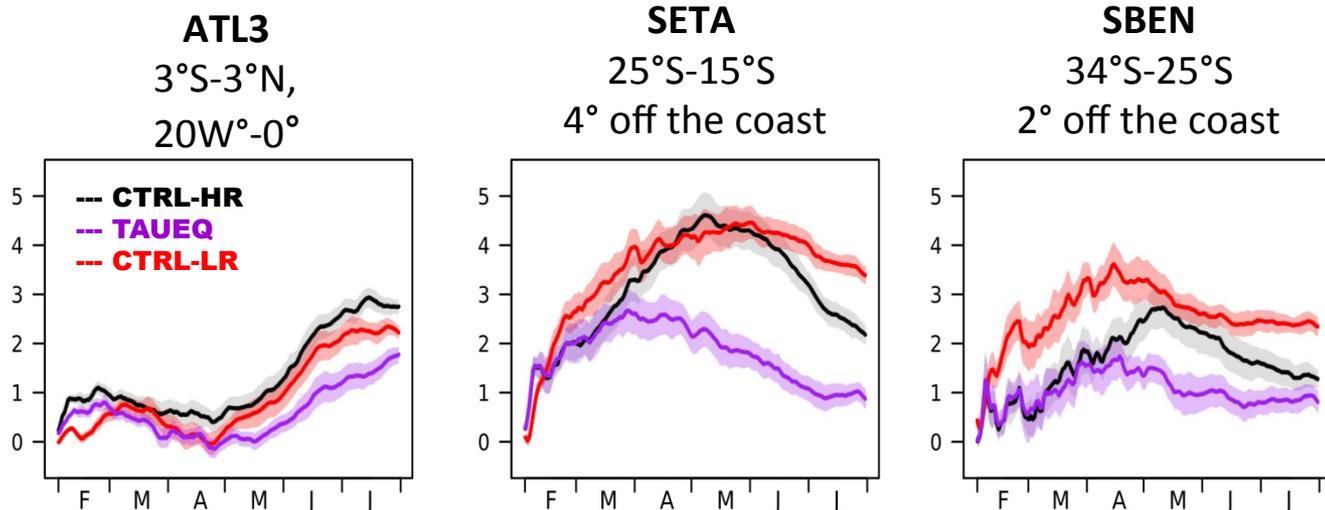
2-months averaged evolution of the SST bias with respect to GLORYS2v3

- **Further analysis** on processes involved to the bias development in CTRL-HR, TAUEQ and CTRL-LR
- **Impact of model resolution**

Analysis strategy



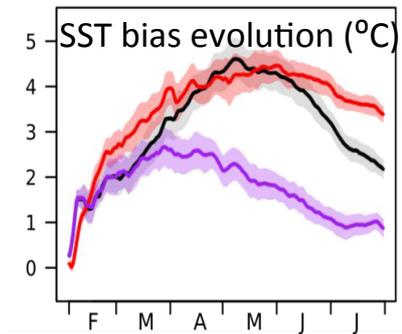
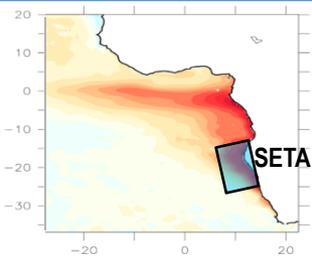
SST bias daily evolution over 3 considered regions (°C):



TAUEQ: improvement in the costal regions by ~50% (from mid-March) and in ATL3 by ~30%

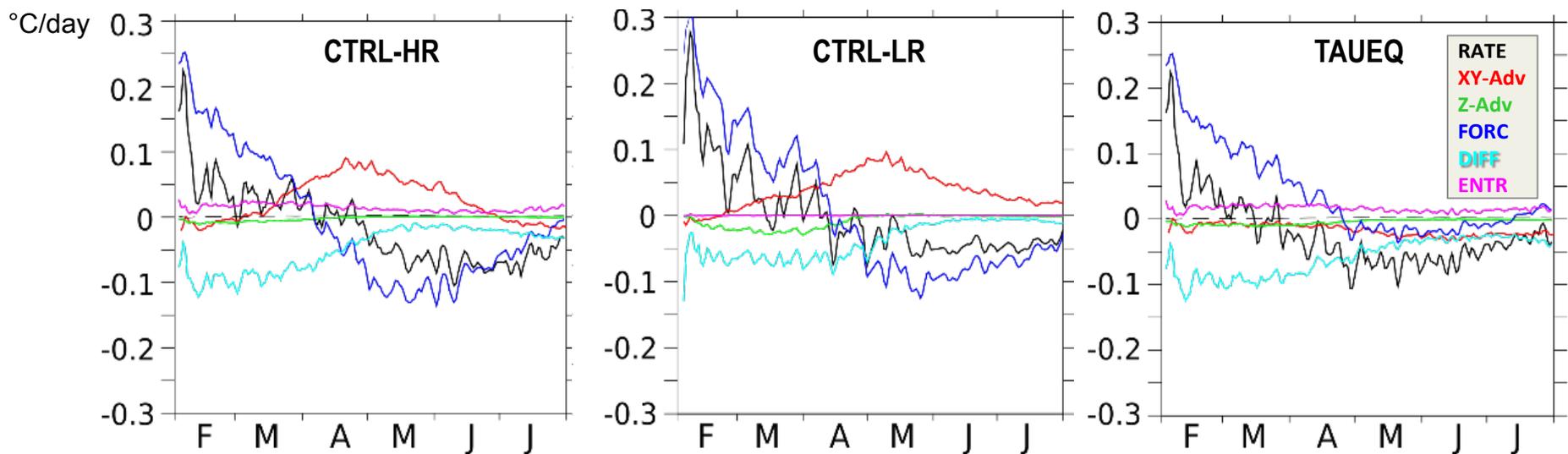
LR: significantly stronger bias in SBEN and in SETA in June-July

SETA region



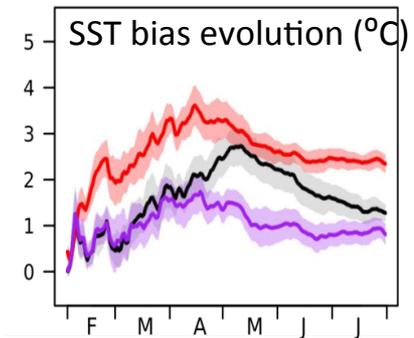
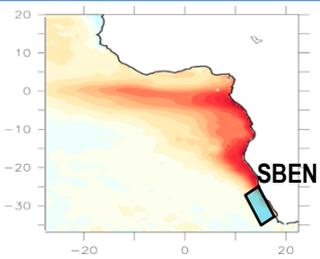
□ Mixed-layer temperature tendency terms

$$\text{TemperRATE} = \text{XY_Adv} + \text{Z_Adv} + \text{Atm.FORC} + \text{Vert.DIFF} + \text{ENTR} + \text{Res}$$



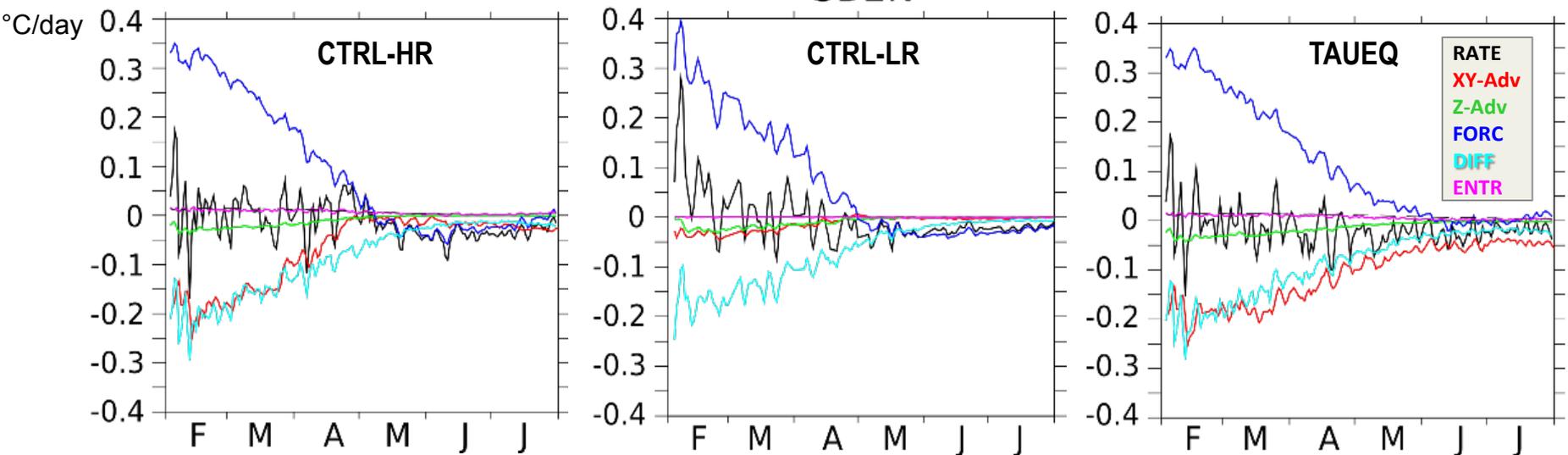
LR: in general similar processes involved, in particular warm horizontal advection
TAUEQ: warm horizontal advection disappears

SBEN region



□ Mixed-layer temperature tendency terms

$$\text{TemperRATE} = \text{XY_Adv} + \text{Z_Adv} + \text{Atm.FORC} + \text{Vert.DIFF} + \text{ENTR} + \text{Res}$$



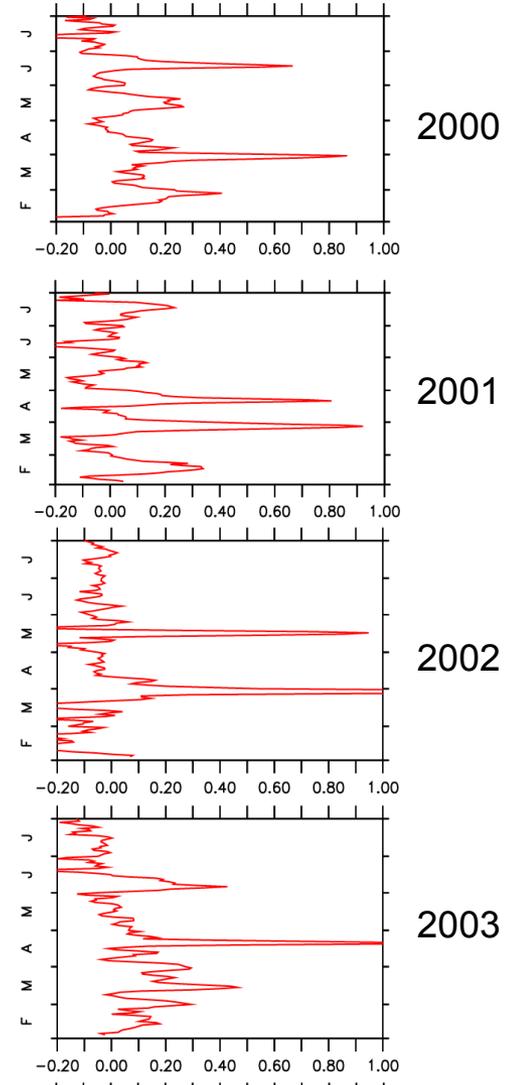
LR: much smaller cold XYadv (Ekman transport/coastal upwelling)

TAUEQ: cold contribution of XYadv(local upwelling) is stronger since there is no “compensation” by warm contribution of XYadv (remote forcing from the Equator)

Warm advection from the Equator



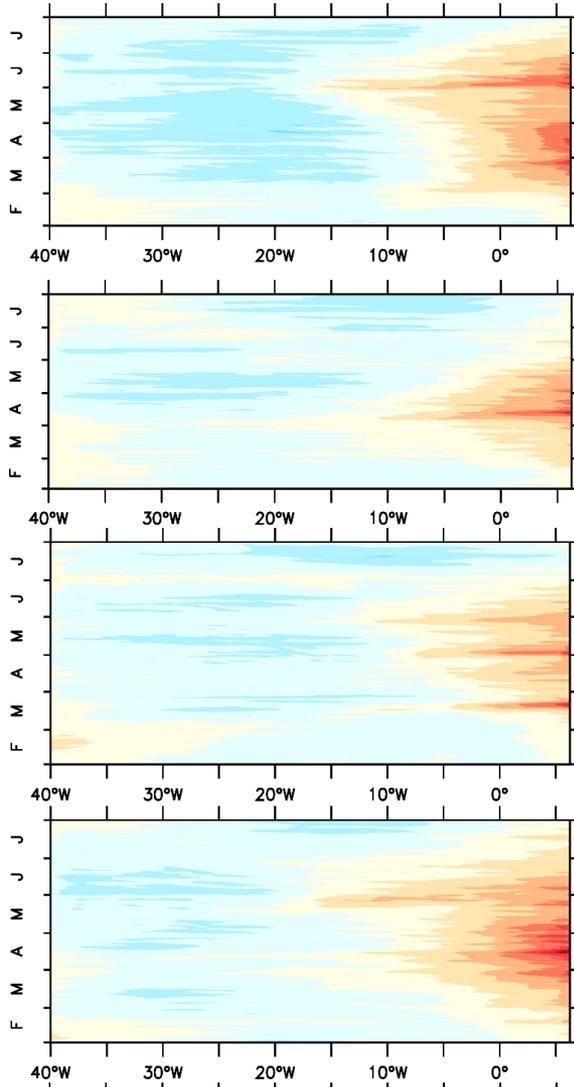
XYadv at 20°S, °C/day
(1st grid cell from the coast)



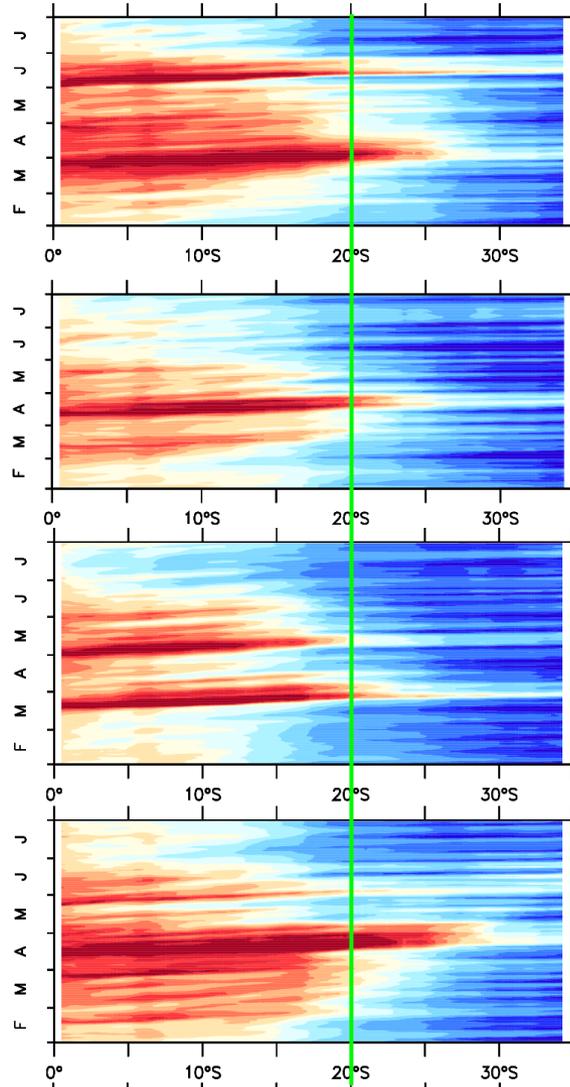
Warm advection from the Equator



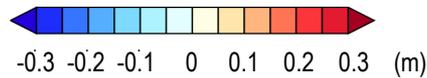
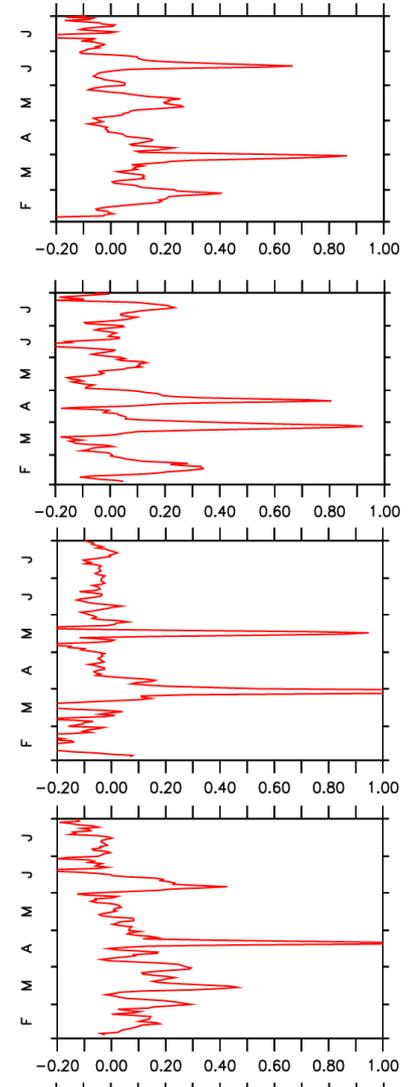
SSH along the equator



SSH along the coast



XYadv at 20°S, °C/day (1st grid cell from the coast)



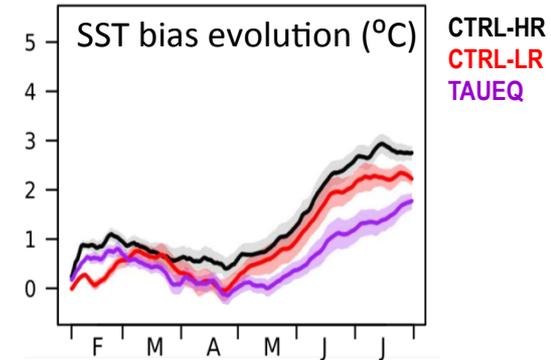
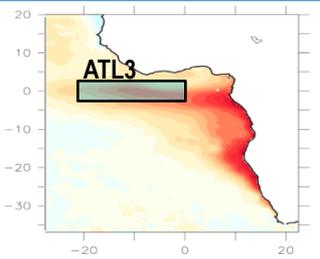
Conclusions

- **Remote forcing from the equator contributes to ~50% of warm SETA bias in CNRM-CM model**
- **This is due to anomalous warm horizontal advection from the Equator that penetrates southward of 25°S and could be associated with propagations of coastal trapped equatorial Kelvin waves**
- **Restoring of the wind stress over the Equator leads to a general improvement of the equatorial mean state (thermocline slope, zonal current) and more realistic representation of the relative contributions of the different terms to the mixed-layer temperature budget in ATL3**
- **Correction of the wind stress over the Equator results in local reducing of the SST bias by ~30% in the equator**
- **In general LR and HR models show very similar evolution of the SST bias and similar associated processes. A significant improvement of the bias in HR is observed locally in the Southern Benguela due to better resolving of fine-scale processes associated with the coastal upwelling.**

Thank you !

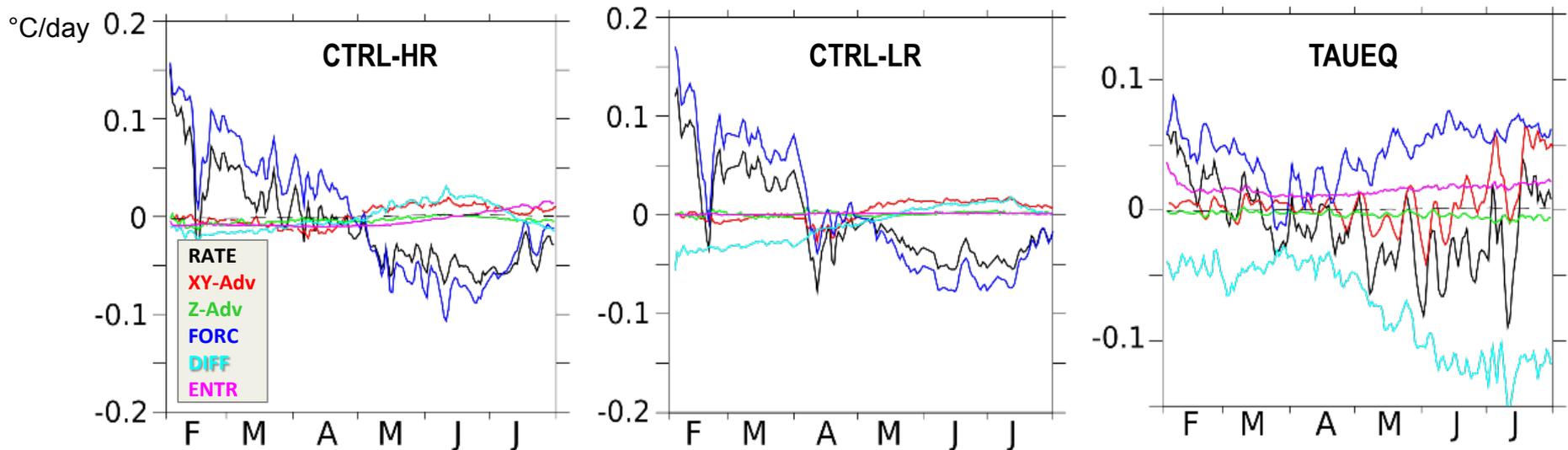
Add-on

ATL3 region



☐ Mixed-layer temperature tendency terms

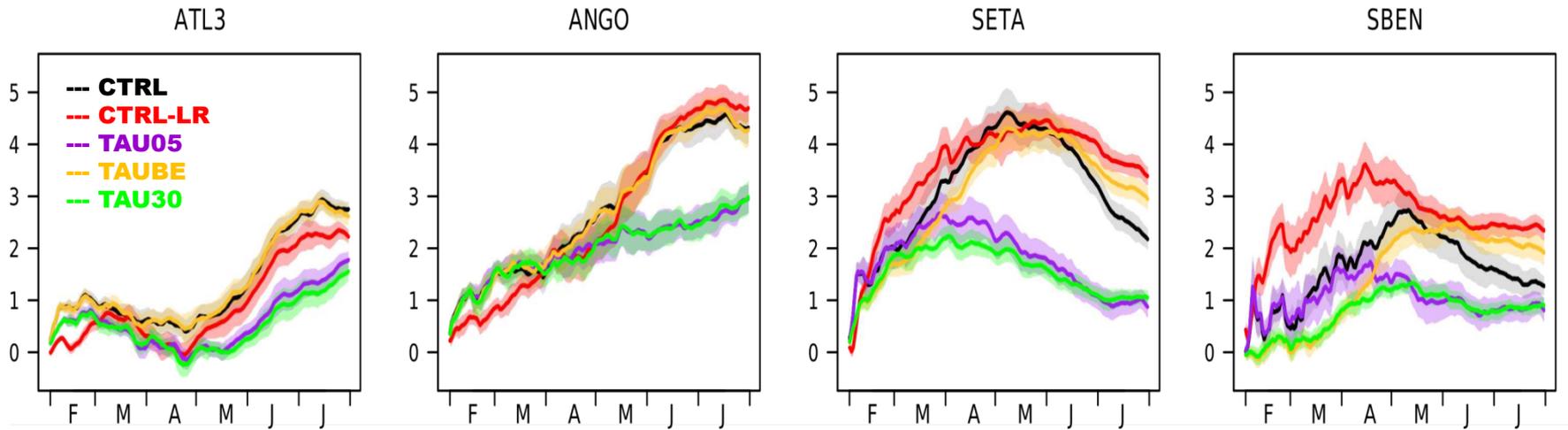
$$\text{TemperRATE} = \text{XY_Adv} + \text{Z_Adv} + \text{Atm.FORC} + \text{Vert.DIFF} + \text{ENTR} + \text{Res}$$



LR: in general similar processes involved

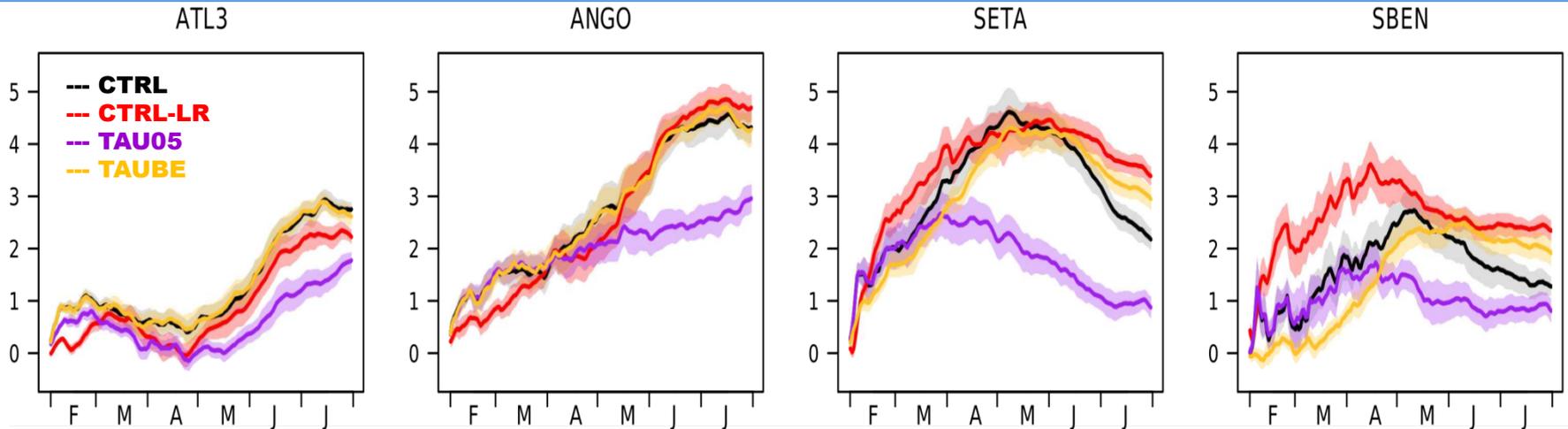
TAUEQ: drastic modification of the ML heat budget, in particular vert.diff and atm.forcing

TAU30: SST bias

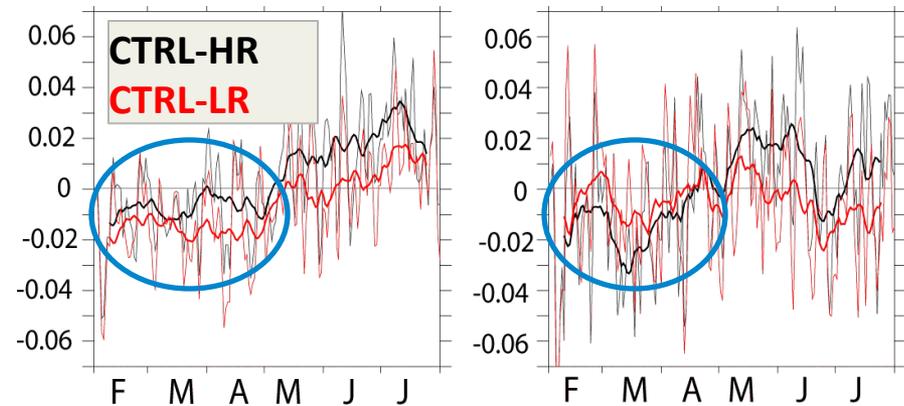


→ Combination of the effects of TAU05 and TAUBE

TAUBE: SST bias

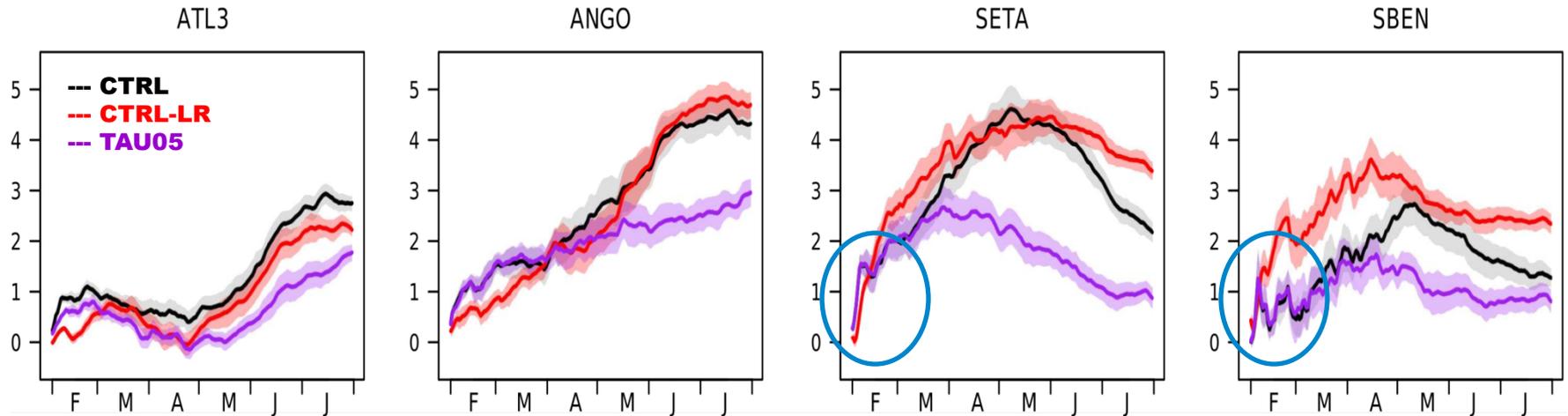


Errors in alongshore wind stress, Pa

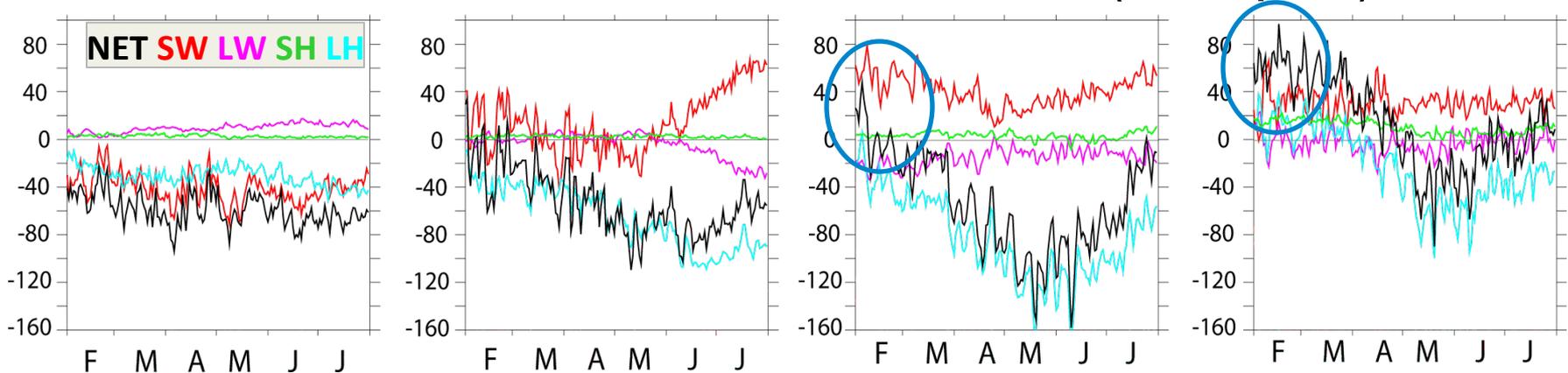


- ✓ Local improving of SST during the first 3 months of hindcasts over SBEN (and in a much lower degree over SETA)
- ✓ Stronger bias in June-July in SBEN and SETA (compensation of errors in CTRL?)
- ✓ There is no remote impact to the Equatorial regions (ATL3 and ANGO) in term of SST

Initial bias: role of heat flux



Errors in net heat fluxes at surface, W/m² (ref: TropFlux)



- ✓ Error in net heat flux at surface may contribute to the rapid initial bias development over SETA and SBEN but it doesn't explain slower initial bias development in ATL3/ANGO